Petrography of Basaltic Rocks

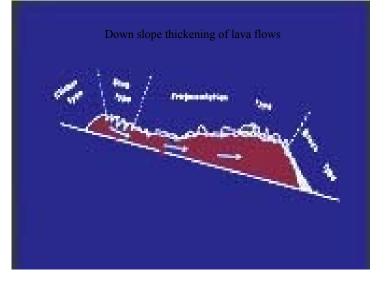
- Fabric
- Classification
- Alteration

Field Relations - Basaltic Rocks

- Intrusions
 - Dikes, sills, plugs, necks
- Extrusions
 - Lava flows (pahoehoe, aa aa)
 - Shield volcanoes
 - Scoria cones
 - Tuff rings
 - Hyaloclastites



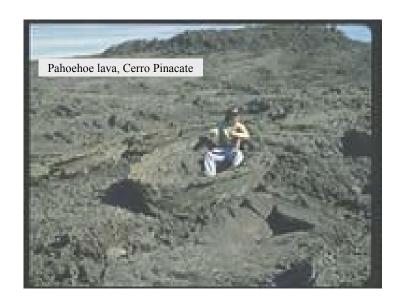




Pahoehoe Flows

- Very fluid lavas
- Smooth surface skin
- Ropy textures, surface pleats
- Flow moves as growing bubbles or buds
- Sometimes a gap at the top of the bud
- Shelly pahoehoe







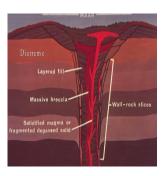
Scoria Cones

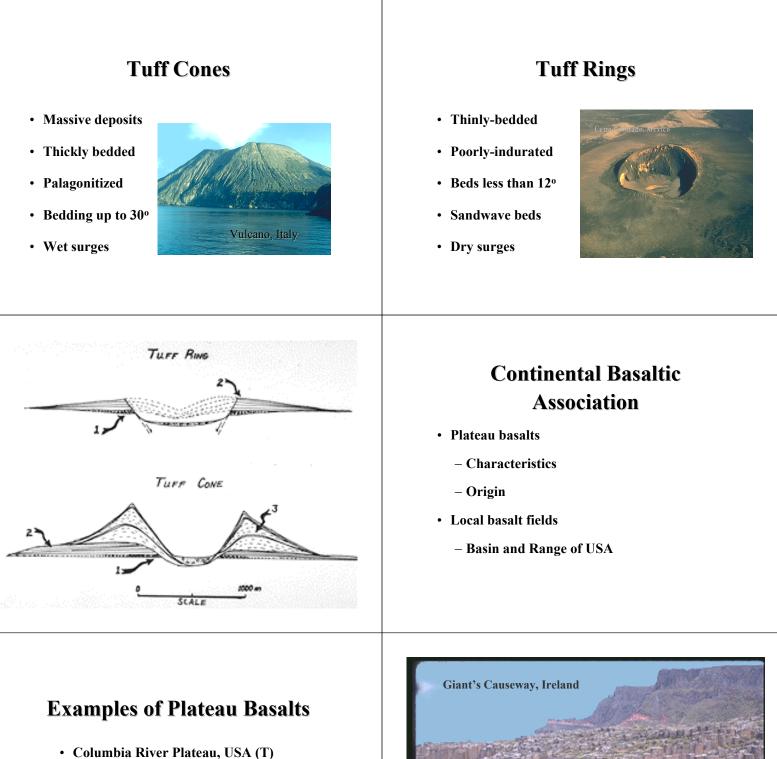
- Simplest and commonest volcanic form
- Characterized by three parameters - Height, width, crater width
- Standard initial slope of 30°
- Conical shape
- Occur in several environments
- McGetchin model of cone growth
- Erosion is systematic



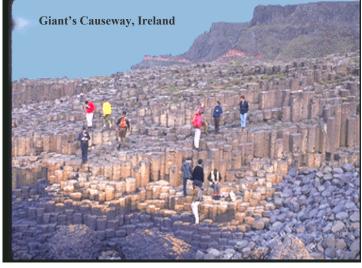
Diatremes

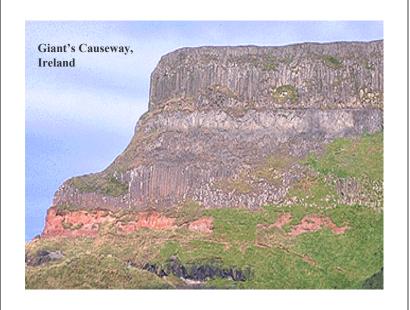
- Breccia pipes
- Kimberlite
- Contains diamonds
- Ultramafic magmas
- Mixture of rocks
- Driven by deep CO₂





- Deccan, India (K-T)
- Parana, Brazil (J-K)
- Keeweenaw, Lake Superior (PreC)
- Karoo, South Africa (J)
- Greenland-Great Britain (K-T)





Plateau Basalt Characteristics

- Fissure eruptions, associated dike systems
- Huge volume (>10⁵ km³)
- Large discharge rate
- May herald the breakup of continents

Chemical Characteristics

- Typically more evolved composition than MORB
 - Higher Si, K, Ti, P, and Ba
 - Lower Mg, and Ni
- Evolved, olivine-poor compositions
 - Suggest some fractionation prior to eruption

Isotopic Evidence

- Low initial Sr isotope ratios (<0.704)
 - Suggest partial melting of upper mantle peridotite
- High initial Sr isotope ratios (>0.704)
 - Suggest contamination with crustal materials

Origin of Plateau Basalts

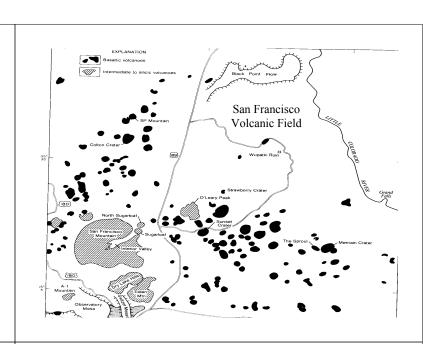
- Low degree of fractionation
- Low initial Sr isotope ratio
- Phase relationships
- Suggest an origin from a peridotite zone within the asthenosphere at a depth between 60 to 100 km

Local Basalt Fields

- May occur in areas of continental extension
 - Basin and Range of western USA
- Characterized by scoria cones and lavas
- Some surround composite andesitic cones
- Minor bimodal basalt-rhyolite (pyroclastic) association

Basaltic Scoria Cone Fields

- 10s to 1000s of cones
- General elliptical shape
- Aspect ratio of 2:1 to 5:1
- 10 to 70 km in length
- Areas of extensional tectonics
- Elongate perpendicular to tension
- Widespread in western USA
- Pinacate example



Small Fields

- North rim of Grand Canyon
- Scoria cones aligned along fault planes



Origin of Local Continental Basalt Fields

- Hot magma from the mantle intrudes rifting crust
- Accumulation of basalt at depth melts silicic crust
- Silicic melt buoyantly rises to shallow chambers
- Shallow chambers erupt to produce evolved pyroclastic deposits

Oceanic Subalkaline Basaltic Association

- Two types of basaltic provinces
 - Intraplate volcanoes (hot spots)
 - Spreading plate boundaries (ocean ridges)
- Iceland
- Oceanic rifts
 - Mid-Atlantic rise
 - East-Pacific rise

Iceland

- Subaerial outcropping of the mid-Atlantic ridge
- No continental sial is present
- Mostly contains quartz tholeiitic
- Minor alkali basalts
- A few eruptive centers
 - Fe-rich andesite, dacite, & rhyoliteProduced by olivine fractionation
- Origin from rising mantle plume?

Icelandic Shields

- Moderate size
- Extremely symmetrical
- Small size >800 m high
- Uniform slope ~ 8°
- Tube-fed pahoehoe lavas



Subglacial Volcanoes

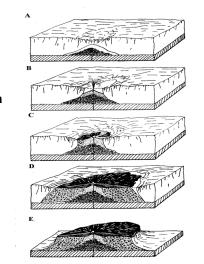
- Pillow lavas
- Pillow breccias
- Hyaloclastites
- Dikes
- Flat top with lava





Sub-glacial

- Sequence of intrusion
- Final form is a table mountain



Oceanic Rifts

- Their lavas comprise 70% of the earth's surface
- Sea floor spreading is the mechanism of their origin

Oceanic Lithosphere

- Layer 1
 - 0 to 1 km thick, sediment
- Layer 2
 - 1 to 3 km thick, basalt flows, pillows breccia, dikes
- Layer 3
 - 4 to 8 km thick, fractured mafic intrusions
- Below layer 3 is is subcrustal peridotite

Ocean Floor Basalts

- MORB
 - Reference composition to other basalt types
 - See book Table 5-5 for chemical characteristics
- Low K₂O content & large-ion lithophile elements
- Originate in the mantle
- Partial melts within the asthenosphere
- Olivine tholeiitic composition (Ol and Hy in norm)

Ocean Floor Lavas

- Evidence of disequilibrium
 - Corroded phenocrysts of Mg olivine and Ca plagioclase
 - Chemically evolved groundmass
 - Anomalous melt inclusions
- Uniform composition of lavas
 - Suggest recurrent mixing in shallow chambers under rifts

Depleted Magma Source

- Several lines of evidence
 - Extremely low concentrations of incompatible elements
 - Rb/Sr ratio too low to yield Sr isotopic ratio
 (~0.703)

Models for Ocean Floor Lavas

- Thin lid model
 - Primitive lavas fed from center of chamber
 - More fractionated materials from margins
- Evolving system
 - Several small chambers at different stages of fractionation
- Strong role of crystal fractionation
 - Supported by presence of mafic cumulate horizons

Ophiolites

- Alpine ultramafic bodies
- Hartzburgitic type
 - Mainly hartzburgite and dunite
 - Minor dikes & veins of other types
 - Can not be the source of basaltic magmas by melting
- Lherzolitic type
 - Mainly lherzolite, minor pyroxenite
 - May yield basaltic magmas by partial melting

Alpine Ultramafic Association

- Steinmann trinity
 - Ultramafic rocks
 - Pillow basalts (spilitic = metasomatized basalt)
 - Chert (with argillite and limestone)
- Origin by obduction
 - Ocean floor thrust onto continental crust during mountain building

Ophiolite Sequence

- Refractory residue of upper mantle hartzburgite
 - Deformed and drained of low-melting point materials
- Overlying fossil magma chambers
- Capping of fractionated basaltic lavas and dikes
 - Sheeted dike complexes