

## Constitution of Magmas

- Hot molten rock
- $T = 700 - 1200$  degrees C
- Composed of ions or complexes
- Phase
  - Homogeneous
  - Separable part of the system
  - With an interface

## Composition

- Most components
  - Low vapor pressure
  - Designated by mole fraction ( $X_i$ )
- Volatile components
  - Mainly exist as a gas
  - Designated by vapor pressure ( $p_i$ )
- Fluid pressure = sum of partial pressures

## Gas law

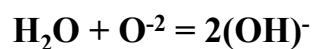
$$PV = nRT$$

## Atomic Structure of Magma

- Quenched to form a glass
- Si & Al are polymerized with O
- Forming networks of Si-O chains
- Short-range structural order

## Structural Model

- Network formers
  - Si, Al
- Network modifiers
  - Ca, Mg, etc
- Dissolved water has a strong effect



## Magma Generation

- Magmas form at perturbations in P,T,X
- Convergent plates
- Divergent plates
- Peridotite mantle source

## **Source Regions**

- Must originate in the mantle or crust
- At Hawaii 60 km deep
- Only 1 to 3% melt in peridotite

## **Melting**

- **Heat of fusion**
  - About 300 times the rock's specific heat
  - Melting of rock consumes much heat
- **Mechanisms for melting**
  - Temperature increase by mass transfer
  - Decompression
  - Changes in composition reducing melting point

## **Temperature Increase**

- **Mechanical deformation**
  - Friction generates heat
- **Mass transfer of rock**
  - Descending oceanic lithosphere
  - Basaltic underplating of continental crust

## **Decompression**

- **Upwelling mantle**
  - Beneath oceanic or continental rift
- **Adiabatic system**
  - Pressure causes all temperature change

## **Changes in Composition**

- **Increase in water pressure**
- **Lowers the solidus**
- **Subduction zones**
  - Peridotite wedge
  - Over subducting oceanic crust

## **Magma From Solid Rock**

- Basalt & peridotite systems
- Granite systems

## **Basalt & Peridotite**

- **Equilibrium fusion**
  - Solid and liquid remain in equilibrium
  - Continuous but limited composition range
- **Fractional fusion**
  - Liquid is immediately removed from host rock
  - Melts are both oversaturated & undersaturated with respect to Si

## **Influence of Pressure**

- Pressure strongly influences the cotectic
- Partial melts of mantle peridotite are basalts
- At higher pressures partial melts are more silica deficient

## **Role of CO<sub>2</sub>**

- Polymerizes melt
- Contracts olivine field
- Favors silica-poor alkali melts
- Repeated melting episodes favors incompatible element enrichment

## **Role of H<sub>2</sub>O**

- Depolymerizes melt & stabilizes olivine
- Partial melts more silica rich
- Favors tholeiitic basalts

## **Mantle-derived Primary Melts**

- **Wide range of melt compositions possible**
- **Fractional crystallization vs. Partial melting**
- **Primary melt**
  - Segregated from peridotite source rock
  - First crystallized minerals similar to mantle source zone
- **Derivative melt**
  - Modified after leaving the source region

## **Granitic Systems**

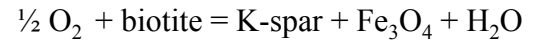
- Impossible to generate granites by partial melting of mantle peridotite or subducted oceanic floor basalt
- Their origin is related to older sialic crust
- Granites concentrated along old subduction zones

## Water Saturation

- Saturated granite melts have 10 to 15% H<sub>2</sub>O
- Natural granite melts have about 4% H<sub>2</sub>O

## Water Undersaturation

- Common granite mineral assemblage  
– Biotite, K-spar, Fe-Ti oxide



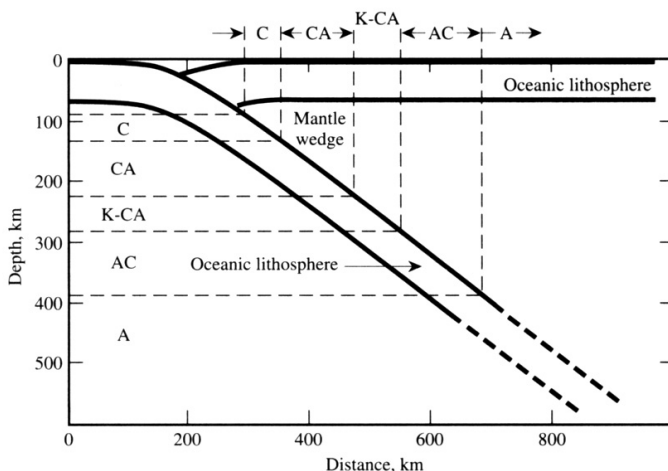
- Excess water drives this reaction to the left
- Hence, most granites are not water saturated

## Origin of Granites

- Partial Melting of lower crust
- Source in mica-amphibolites
- Contain 1-2% H<sub>2</sub>O
- Lowest T melts are K-rich granite
- Higher T, deeper melts are Ca-rich granodiorite

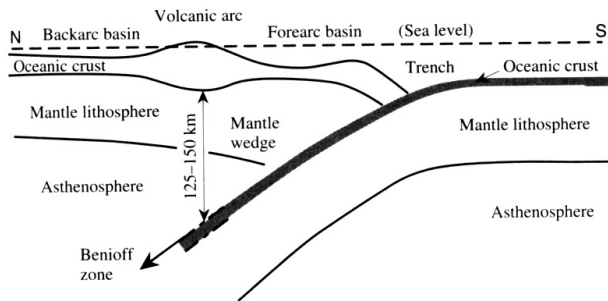
## Subduction Zone Magma

- **Subducted slab**  
– Mafic primary melts
- **Peridotite mantle wedge**  
– Mafic primary melts



## Dehydration Beneath Orogen

- **Large amount of water in oceanic slab**  
– Water in pore space  
– Water in alteration minerals
- **Heating dehydrates the slab**
- **Liberated water promotes partial melting of peridotite**
- **Composition is Si-saturated tholeiite**



## Magma Diversification

- Magmatic differentiation
- Gravitational settling
- Liquid immiscibility

## Crystal-liquid Fractionation

- Regular pattern of compositional variation
- Variation of MgO is a good measure of olivine fractionation
- Computer mixing programs can be used

## Magma Mixing

- Two different magmas may blend to produce a hybrid
- Common with calc-alkali magma
- Blended magmas should have linear composition with the parents